Dissertation Methodology

Notes:

* Mention that this method is pretty much the same as Uieda only difference is the cross validation.
* Brief overview on parametrization, forward problem, inverse problem, regularisation, and Bott’s method.
* Cross validation – split up have training size ranging between 2/3, 3/4, and 4/5 to see the mean error between gravity and seismic estimates. Want to see how many points are needed to still get relatively accurate values (small MSE’s) so we know that when seismic point estimates are not present the gravity moho calculated is still accurate (at around 2.3-2.5km error).
* Statement about using python software.

Methodology:

To accurately obtain a model of the Moho depth one must first remove all other effects that contribute overall gravitational values recorded over an area. This is achieved through correcting for things such as removing the scalar gravity of an ellipsoidal reference Earth (the Normal Earth) and then the removal of all other effects via a Bouguer correction otherwise known as the correction for topography. Initially though the effect of the Normal Earth needs to be removed from the same point as where the gravity observation was made and is calculated from the closed-form solution in Li & Götze (2001). The value obtained here is called the gravity disturbance and can be seen in the equation below (insert gravity disturbance equation once on .tex file). (Also include Figure 1 in Uieda & Barbosa paper, stages of gravity correction).

The gravity disturbance is still not a direct result of the change in density associated with the Moho discontinuity but also an amalgamation of topography with reference to the normal ellipsoid, variations of density in the crust (e.g. sedimentary basins and igneous intrusions), anomalies below the upper mantle, and mass deficiency due to the oceans. These effects are removed through a topography correction, (insert topography correction equation). This is the method used in Uieda & Barbosa (2017) and assumes that the effects of other crustal and mantle sources are negligible, and after this correction the gravitational values attained are purely a result of the density variations either side of the Moho discontinuity. Seeing as the data is obtained for a sufficiently large area (South America) this correction for topography amongst other things is calculated using tesseroids as part of a spherical Earth approximation. Effects of the tesseroids are calculated using a GLQ integration presented in Asgharzadeh et al. 2007) and improved upon in Uieda et al. 2016 through the adaptive discretization scheme. (insert Figure 2 in Uieda & Barbosa, tesseroids).

Overview of the methodology of calculating a gravity derived Moho model from Uieda (2017)

Note: for each section don’t write more than 200 words, around 1000 overall for this bit.

Parametrization, the forward problem, inverse problem, regularization, Bott’s method, estimating the parameters:

As the implementation of cross validation to estimate uncertainty in the model is implemented as part of the code used in the Uieda & Barbosa (2017) paper the method in calculating the model is largely similar. For context, an overview of this method will be given but for more detail see Uieda & Barbosa (2017).

Upon the calculation of the Bouguer disturbance from the removal of topography, sediments etc. the forward model is parameterized by discretizing the anomalous Moho onto tesseroids. The forward model aims to calculate the difference between the Normal Earth Moho and the true Moho depth, and depending on which is shallower will result in either a positive (red) or negative (blue) density contrast displayed by the colour of the tesseroids. The overall absolute value of the density contrast is predetermined, this produces a nonlinear problem with equation, (insert equation 3 from Uieda, don’t include i values just generic equation). Where d is the data vector, p is the parameter vector containing Moho depths, and f is the non-linear function.

Leading on to the inverse problem the parameter vector is estimated using least squares that reduces the misfit to the data, (insert least-squares equation). Where d0 is the observed gravity data, the equation means that this is a non-linear inverse problem, but we can calculate the minimum error using optimization, where a perturbation vector Δp0 is iterated until a minimum is reach which leads to the minimum value of φ(p).

The optimization of the least squares estimate however is not enough for estimating the relief associated with the Moho and needs regularization in the form of a first-order Tikhonov regularization (Tikhonov & Arsenin 1977) to ensure smoothness in the model and no sharp jumps in Moho depth, to provide a realistic model. (inset regularization equation). R a matrix composed of first order differences between tesseroid depths. This along with the least squares estimate leads to an inverse problem that is solved by minimising the goal function, (insert goal function), µ is the regularization parameter that helps control the fit to the observed data and the smoothness.

After the rearrangement and substitution of equations we arrive at a linear equation system which can calculate the real Earth Moho depths with reference to the Normal Earth Moho. (insert linear equation system equation), where Ak is the Jacobian matrix, and Δpk is the parameter perturbation vector.

(need to write section about Bott’s method and how it combines with the rest of the method)